

1 **Observed trend of precipitation extreme in Kalimantan**

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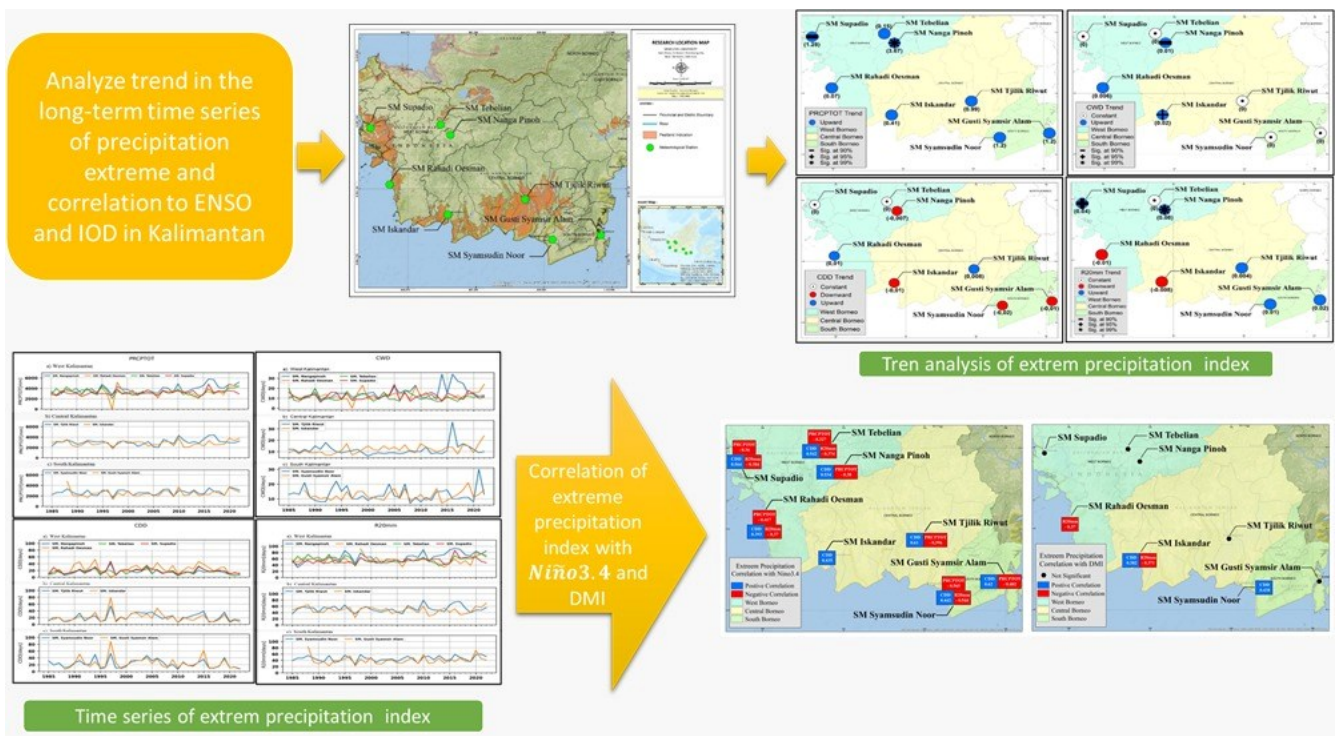
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13 **GRAPHICAL ABSTRACT**



14

15 **Abstract**

16 It has been suggested that the global warming due to an increase in green house gas concentrations  
 17 could lead to extreme climate events, such as precipitation extreme and droughts. This study is  
 18 designed to analyze trend in the long-term time series of precipitation extreme in Kalimantan, focusing  
 19 on observed data obtained from eight meteorological stations spanning from January 1985 to December  
 20 2022. Statistical analyses were employed to discern patterns and trends in extreme precipitation events.  
 21 The findings reveal a discernible long-term trend characterized by fluctuations in precipitation  
 22 extremes over the study period. The trend of total annual precipitation (PRCPTOT) has increased at all  
 23 observation stations. Especially at SM. Nangapino and Supadio, the PRCPTOT trend increased  
 24 significantly at the 99% and 90% levels. Meanwhile, the increase in consecutive wet days (CWD)  
 25 trends only occurred at SM. Nangapino and Rahadi Oesman (West Kalimantan), and SM. Iskandar  
 26 (South Kalimantan). A significant trend was found in SM. Nangapino (90%) and Iskandar (95%). In  
 27 addition to experiencing an increase in the R20mm trend, there was a decrease in the trend. Significant

28 trend increases at the 99% and 90% levels occurred in SM. Nangapinoh and Supadio. In addition, an  
29 increase in CDD trend occurred only in SM. Rahadi Oesman and Tjilik Riwut, although not significant.  
30 Meanwhile, the majority of extreme precipitation indices have a greater correlation with ENSO than  
31 the IOD.

32  
33 *Key words: climate extreme; Kalimantan; Mann-Kendall test; precipitation extreme*

## 34 35 **1. Introduction**

36 Precipitation extremes play a pivotal role in shaping the hydrological and ecological dynamics of  
37 regions worldwide, exerting significant impacts on water resources, agriculture, infrastructure,  
38 ecosystems, and soil erosions (Yang et al., 2016; Fei et al., 2021; Kastridis et al., 2024; Piacentini et al.,  
39 2018). Kalimantan, the Indonesian portion of the island of Borneo, is no exception to the influence of  
40 such extremes. Situated within the equatorial belt, Kalimantan experiences a tropical climate  
41 characterized by high precipitation variability, with precipitation patterns profoundly influencing the  
42 region's socioeconomic and environmental landscapes (Estiningtyas et al., 2024). Understanding the  
43 trends and patterns of precipitation extremes in Kalimantan is therefore crucial for effective water  
44 resource management, disaster preparedness, and sustainable development initiatives.

45 Over recent decades, increasing attention has been devoted to the study of climate variability and  
46 its implications for regions globally (IPCC, 2023). In Kalimantan, where precipitation extremes can  
47 trigger floods, landslides, and other natural disasters, the need for comprehensive analyses of observed  
48 trends in extreme precipitation events is particularly pressing. Such analyses provide valuable insights  
49 into the changing climatic conditions and their potential consequences for the region's inhabitants and  
50 ecosystems. Despite the significance of understanding precipitation extremes, relatively few studies  
51 have focused explicitly on this topic in the context of Kalimantan. Existing research often lacks a

52 comprehensive examination of long-term trends, instead focusing on short-term analyses or specific  
53 events.

54 Precipitation trend analysis in the West Kalimantan for the Kubu Raya and Mempawah districts  
55 was conducted from 2000-2019. The results show an increase of monthly maximum consecutive 5-day  
56 precipitation (RX5day), monthly maximum 1-day precipitation (RX1day), and number of days when  
57 precipitation above 50mm (R50mm). Meanwhile, the consecutive dry days (CDD) experienced a  
58 decreasing trend (Aditya et al., 2021). Another study in the South Kalimantan revealed an increasing in  
59 precipitation of about 25mm per year over a period of 2000 - 2020 (Sukmara et al., 2022). Recent study  
60 based on Integrated Multi-Satellite Retrievals for GPM (IMERG) version 6 data over a period of 2001  
61 – 2020 shows an increasing trend in precipitation extreme indices in new capital city of Indonesia in  
62 the east Kalimantan (Marzuki et al, 2023). The precipitation frequency-based indices, namely R20mm  
63 and R50mm, indicate an increasing trend in the last 2 decades. Similarly, the R5Xday index also  
64 revealed an increasing trend. On the other hand, the total amount of precipitation, the RX1day as well  
65 as the SDII (annual total precipitation divided by the number of wet days in the year) has shown a  
66 decreasing trend over the last 2 decades.

67 It should be note that the spatio-temporal variability of precipitation in the Indonesian region is  
68 influenced by the local and regional air-sea interaction (Aldrian and Susanto, 2023). On seasonal time  
69 scale, the monsoon system mostly drives the seasonal variation of precipitation over the Indonesian  
70 region (Mulsandi et al., 2024). Most of the Indonesian region experiences wet season during the  
71 northwest monsoon, while during the southeast monsoon Indonesia experiences a dry season. On  
72 interannual timescale, the Indo-Pacific climate modes, namely the El Niño-Southern Oscillation  
73 (ENSO) and the Indian Ocean Dipole (IOD), modulate the precipitation over the Indonesian region  
74 (As-Syakur et al., 2013; Lestari et al; 2018; Kurniadi, 2021). El Niño and/or positive IOD conditions  
75 often result in decreased precipitation and increased temperatures, creating dry and drought-prone  
76 conditions, particularly in the archipelago's forested regions. These dry conditions significantly elevate

77 the risk of forest fires in Indonesia (Iskandar et al., 2022; Nurdiati et al., 2021), which are exacerbated  
78 by the accumulation of flammable biomass in peatlands and forests. The combination of reduced  
79 precipitation, higher temperatures, and prolonged dry spells during El Niño events creates ideal  
80 conditions for the ignition and spread of fires (Hooijer et al., 2006; Wooster et al., 2012). These fires  
81 not only pose a threat to human health and safety but also cause extensive damage to ecosystems,  
82 biodiversity, and the economy, with impacts ranging from loss of habitat for endangered species to  
83 disruptions in agriculture and tourism (Sarmiasih & Pratama, 2019; Lohberger et al., 2017).

84 Nevertheless, there is still a gap in the literature regarding the observed trends of precipitation  
85 extremes in Kalimantan over an extended period. Addressing this gap is essential for enhancing our  
86 understanding of the region's climate dynamics and improving the accuracy of future climate  
87 projections and risk assessments. Therefore, this study aims to bridge this gap by providing a  
88 systematic analysis of the observed trend of precipitation extremes in Kalimantan. Drawing upon  
89 observed meteorological datasets spanning multiple decades (1985 – 2022), this study employed  
90 statistical analysis to quantify changes in extreme precipitation events over time. By identifying trends  
91 and patterns in precipitation extremes, we seek to elucidate the underlying drivers and mechanisms  
92 driving these changes link to ENSO and/or IOD events. Additionally, we explore the potential  
93 implications of observed trends for forest fire management, and climate adaptation efforts in  
94 Kalimantan.

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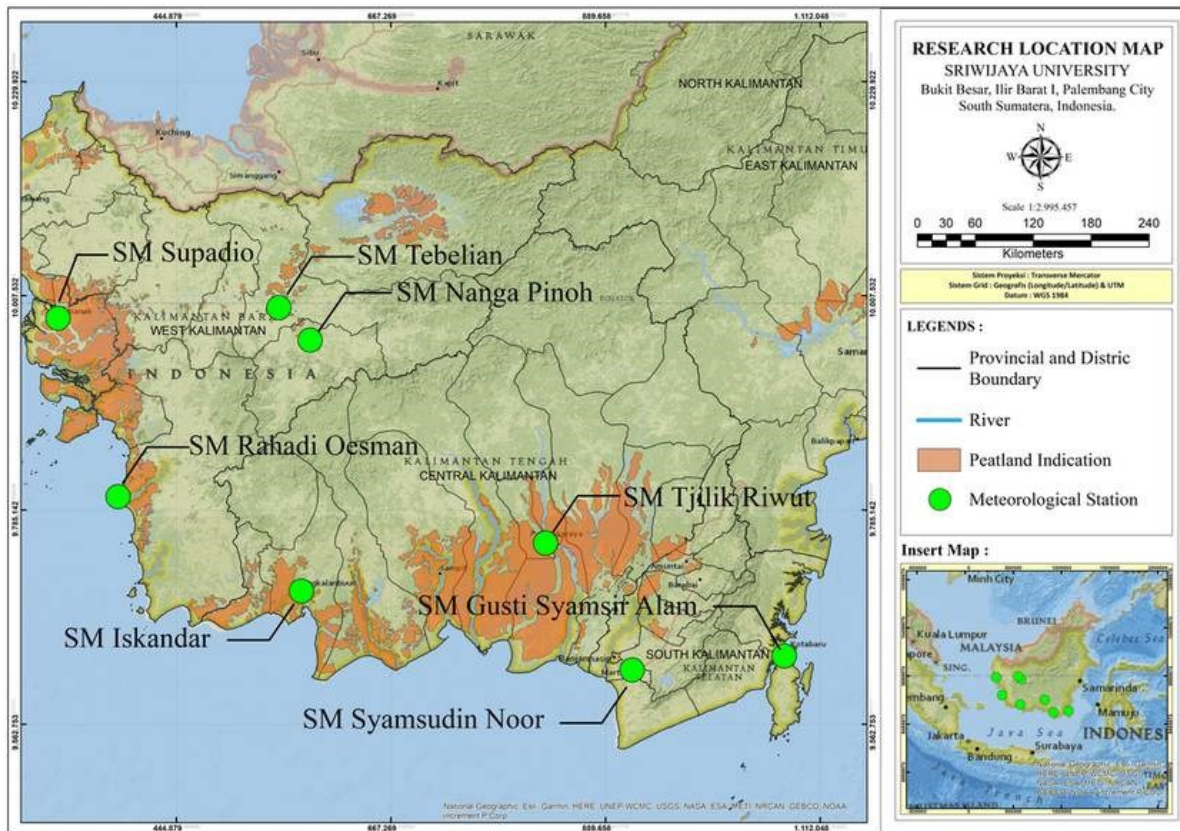
## 96 **2. Materials and Methods**

### 97 **2.1. Study Area**

98 Kalimantan is the third largest island in the world with, an area of 746,000 km<sup>2</sup> (Wooster et al.,  
99 2012). It is located at coordinates 7°N - 4.5°S, 108°E - 119°E, crossed by the equator. Borneo Island is  
100 mainly included in Indonesia's territory, partly in Malaysia and a small part in Brunei Darussalam. The

101 Indonesian part is divided into five provinces: West, Central, East, North, and South Kalimantan  
102 (Yulianti & Hayasaka, 2013). Almost the entire region is covered by tropical rainforests and peat.  
103 Forest types in Kalimantan include mangrove forests, peat swamps, freshwater swamp forests, mixed  
104 dipterocarp forests, montane forests, forests on limestone and ultrabasic soils (Area et al., 2007). The  
105 world's tallest dipterocarp forests are found in Borneo. These forests are vulnerable to drought and fire  
106 (E. Guhardja, et.al, 2000). The forest area in Kalimantan reaches 30% of the island area, which is about  
107 24 million hectares, so Kalimantan has a lot of endemic fauna and flora.

108 Kalimantan is also the most biodiverse region in the world. Of the 15,000 plant species, 6,000  
109 are endemic (Area et al., 2007). The fauna includes 268 species of mammals, 523 species of birds, 147  
110 species of amphibians, 227 species of reptiles, and 738 species of freshwater fish (Widjaja, 2014).  
111 Kalimantan is one of the world's places with orangutans, elephants, and rhinos living side by side.  
112 Besides being rich in flora and fauna, Kalimantan has 20 large rivers in the highland area. The forests  
113 in Kalimantan are inhabited by around 4 million Dayak tribes (the original occupants), divided into  
114 hundreds of different ethnic groups with their respective languages and cultures. Therefore, it is  
115 essential to protect the forest for the community and ensure water and food supply. In reality,  
116 Kalimantan has been transformed into agricultural land and plantations. These fragmented landscapes  
117 have been acquired through logging processes and human-caused burning events. If this process  
118 coincides with El Nino and positive IOD phenomena, then fires can spread to undisturbed forests and  
119 peatlands (Langner et al., 2007; Goldammer, 2007; E. Guhardja et.al, 2000 ). Therefore, this research  
120 was conducted in peat areas located in West, Central, and South Kalimantan, as shown in Figure 1.



121

122 Figure 1. Research locations covering West Kalimantan, Central Kalimantan, and South Kalimantan-  
 123 Peat areas are marked with brown color.

124

125 **2.2 Materials**

126 This study uses precipitation data from the Meteorology, Climatology, and Geophysics Agency  
 127 (BMKG) through the website at <http://dataonline.bmkg.go.id>. There are 14 stations in West  
 128 Kalimantan, five in Central Kalimantan, and three in South Kalimantan. However, only eight stations  
 129 from the three provinces have complete data for 38 years. The data used is from January 1985 to  
 130 December 2022, recorded at stations as shown in Table 1.

131

Table 1. Location of BMKG Stations used in the Study

No	Station Name	Province	Coordinates	
			Latitude	Longitude
1	SM. Rahadi Oesman	West Kalimantan	-1.80000	109.97000
2	SM. Nangapinoh	West Kalimantan	-0.42000	111.47000
3	SM. Supadio	West Kalimantan	-0.14206	109.45000

4	SM. Tebelian	West Kalimantan	0.06000	111.47000
5	SM. Tjilik Riwut	Central Kalimantan	-2.22000	113.95000
6	SM. Iskandar	Central Kalimantan	-2.73000	111.66000
7	SM. Syamsudin Noor	South Kalimantan	-3.44200	114.75400
8	SM. Gusti Syamsir Alam	South Kalimantan	-3.30000	116.17000

132 SM: meteorological station

133

## 134 2.3 Methods

135 Four indices from the Expert Team on Climate Change Detection and Indices (ETCCDI) were  
 136 selected to characterize extreme precipitation. The indices are PRCPTOT (total annual precipitation),  
 137 CDD (number of consecutive dry days), CWD (number of consecutive wet days), and R20mm (number  
 138 of days with at least 20mm of precipitation) (Zhang et al., 2011).

### 139 2.3.1 Man-Kendall Test

140 The Mann-Kendall test is a non-parametric test used to determine the data trend based on the  
 141 relative ranking of a specific period (Mann, 1945). This test does not have to fulfill the assumption of  
 142 normality. The Mann-Kendall test can be calculated using the following equation (Kendall, 1948):

$$143 \quad S = \sum_{k=1}^{n-1} \sum_{j=k+1}^n \text{sgn}(x_j - x_k), \text{sgn}(x_j - x_k) = \begin{cases} +1, & \text{if } (x_j - x_k) > 0 \\ 0, & \text{if } (x_j - x_k) = 0 \\ -1, & \text{if } (x_j - x_k) < 0 \end{cases} \quad (1)$$

144 A positive S value indicates an increasing trend, but a negative S value indicates a decreasing trend.

145 The variance of the S value can be calculated using the following equation:

$$146 \quad \text{Var}(S) = \frac{n(n-1)(2n+5)}{18} \quad (2)$$

147 If n is greater than 8, S is statistically close to a normal distribution. Statistical tests are performed  
 148 using the typical distribution approach, and the standard Z test statistic is calculated using the following  
 149 equation:



$$Z = \begin{cases} \frac{S-1}{\sqrt{\text{var}(S)}}, & \text{if } S > 0, \\ 0, & \text{if } S = 0, \\ \frac{S+1}{\sqrt{\text{var}(S)}}, & \text{if } S < 0. \end{cases} \quad (3)$$

A positive Z value indicates an increasing trend, while a negative Z value indicates a decreasing trend.

### 2.3.2 Sen's Slope Estimator

The slope of Sen provides information on how much the average extreme precipitation changes per year. This trend test is assumed to be linear and quantifies time-varying data. Sen's slope test is better than the linear regression test because the test is not affected by outliers and data errors (Sen, 2013). The equation used is as follows:

$$Q_i = \frac{(x_j - x_i)}{j - i}, i = 1, 2, 3, \dots, N, \quad (4)$$

$x_j$  and  $x_i$  data at time  $j$  and  $i$  respectively Where  $j > i$ . The  $N$  values of  $Q_i$  are sorted from smallest to largest, then Sen's Slope uses the median  $Q_i$  ( $Q_{med}$ ) calculated by the equation below:

$$Q_{med} = \begin{cases} Q_{\lfloor \frac{N+1}{2} \rfloor}, & \text{if } N = \text{odd}, \\ \frac{Q_{\lfloor \frac{N}{2} \rfloor} + Q_{\lfloor \frac{N+2}{2} \rfloor}}{2}, & \text{if } N = \text{even}. \end{cases} \quad (5)$$

The turning point of each extreme rainfall index trend can be determined using the Sequential Mann Kendall Test (SqMK). Detailed information about SqMK can be found at (Bisai et al., 2014; Stathi et al., 2023).

## 2.4 Partial Correlation

Partial Correlation determines the influence of ENSO and IOD on the extreme rainfall index. This correlation corresponds to more than one bivariate relationship and is distinguished using subscripts. The coefficient of partial correlation is symbolized by  $r_{y.x.z}$ , where the variable to the right

168 of the point is the control variable. So,  $r_{yx.z}$  this means a partial correlation coefficient that measures  
169 the relationship between variables x and y while controlling for variable z. The equation used to  
170 calculate the partial correlation coefficient is (Healey, 2012; Iskandar et al., 2013):

$$171 \quad r_{yx.z} = \frac{r_{yx} - (r_{yz})(r_{xz})}{\sqrt{1 - r_{yz}^2} \sqrt{1 - r_{xz}^2}} \quad (6)$$

## 172 **3. Results**

### 173 **3.1. Annual precipitation analysis**

174 The spatial distribution of annual precipitation in West, Central, and South Kalimantan can be  
175 seen in Figure 2. The lowest annual precipitation total for West Kalimantan is in SM. Rahadi Oesman,  
176 which amounted to 2,041.3 mm/year, occurred in 2014. At the same time, the highest is in SM.  
177 Nangapinoh at 5,781.7 mm/year, which occurred in 2016. In the Central Kalimantan region, SM.  
178 Iskandar has the lowest total annual precipitation. Iskandar, which occurred in 1997, amounting to  
179 1,702.3 mm/year. At the same time, the highest is in SM. Tjilik Riwut at 4,405.5 mm/year, which  
180 occurred in 2017. South Kalimantan has the lowest total annual precipitation in SM. Gusti Syamsir  
181 Alam at 1309 mm/year, which occurred in 1997. At the same time, the highest total annual precipitation  
182 is in SM. Gusti Syamsir Alam of 4780 mm/year, which occurred in 1988. The variation of Kalimantan's  
183 total annual precipitation ranges from 1300-5800 mm/year. The lowest total annual precipitation is in  
184 South Kalimantan, while the highest is in West Kalimantan. The high precipitation variability in  
185 Kalimantan indicates that Kalimantan is experiencing the impact of climate change.

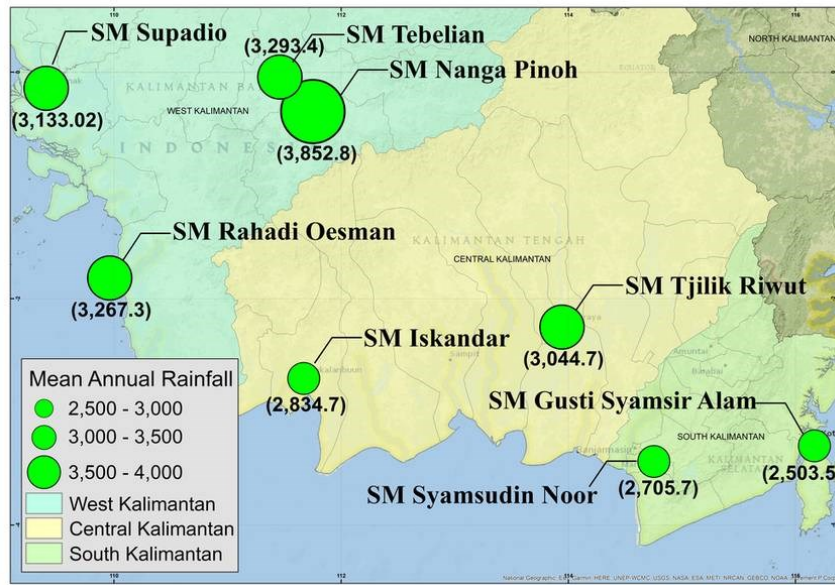
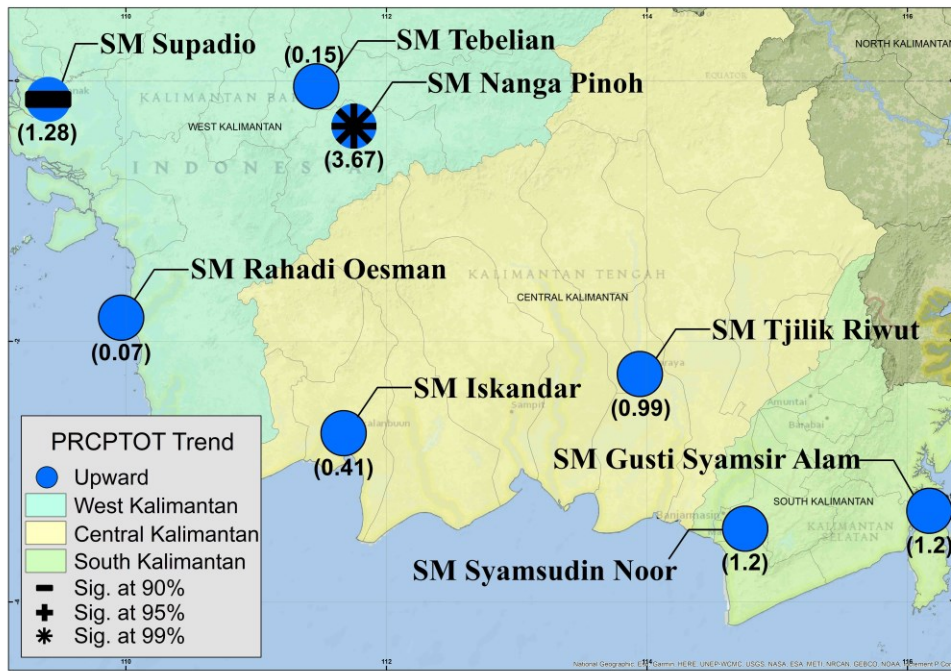


Figure 2. Total Annual Precipitation

### 3.2. Trend analysis of extreme precipitation index

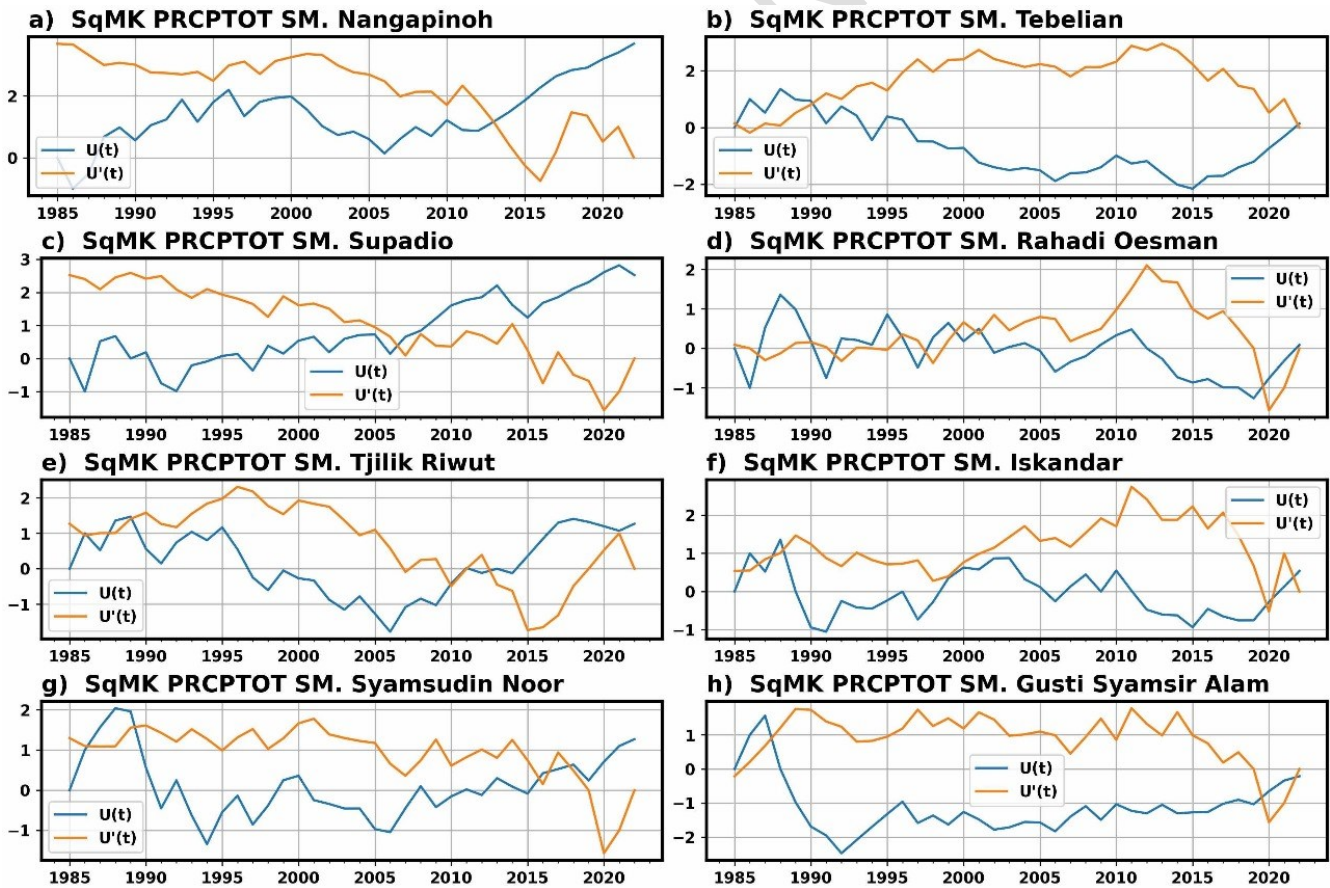
Figure 3 shows the trend of PRCPTOT at the eight Meteorological Stations. It can be seen that the trend of PRCPTOT at eight stations has increased. A significant increase occurred at SM. Nangapinoh (3.67 mm/year) and SM. Supadio (1.27 mm/year) with significance levels of 99% and 90%, respectively. Figure 4 shows that changes appearing in the PRCPTOT time series at Nangapinoh SM started in 2013, while SM Supadio started in 2007. The maximum trend occurred in SM. Nangpinoh, experienced an increase in PRCPTOT per year of 3.67mm. At the same time, the smallest trend increase occurred in SM. Rahadi Oesman by 0.07 mm/year. The increasing trend in eight stations shows that West, Central, and South Kalimantan tend to be in wet conditions.



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Figure 3. PRCPTOT Trends in West Kalimantan, Central Kalimantan and South Kalimantan

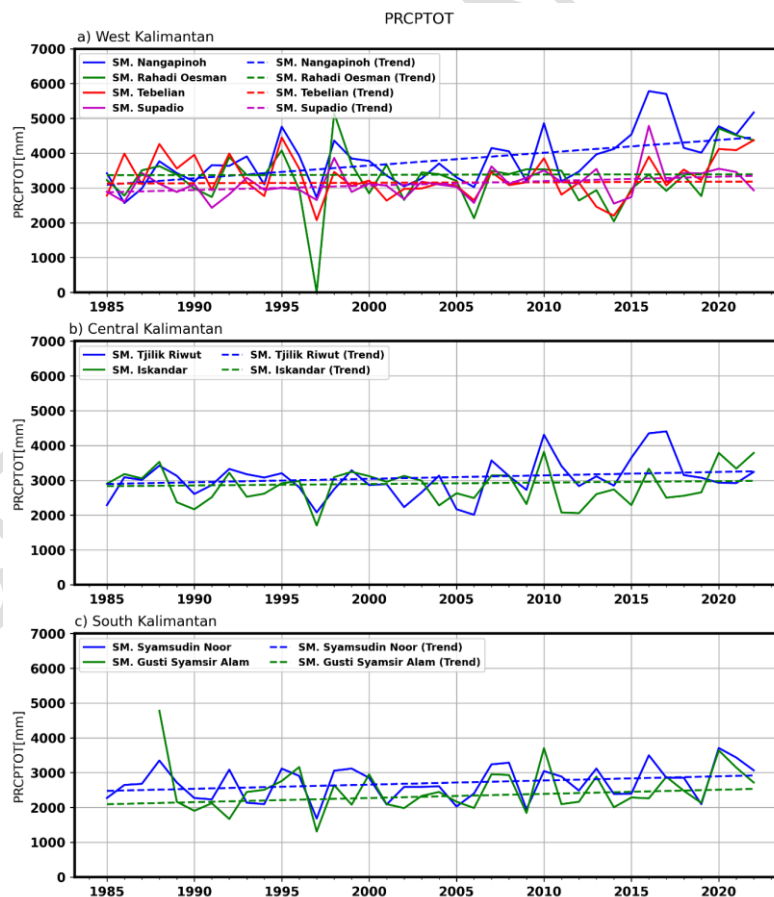


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Figure 4. Graphical representation of the Mann–Kendall trend test of the PRCPTOT time series

202 The variation of PRCPTOT over 38 years in the three regions can be seen in Figure 5. The  
 203 maximum PRCPTOT in West Kalimantan is found in SM. Nangapinoth of 5,781.7 mm which occurred  
 204 in 2016 (Figure 5a). At the same time, the minimum PRCPTOT is found in SM. Rahadi Oesman of  
 205 2,041.33 mm, which occurred in 2014. For the Central Kalimantan region, the variation of PRCPTOT  
 206 can be seen in Figure 5(a). Figure 5(b) shows the maximum PRCPTOT value is found at SM. Tjilik  
 207 Riwut amounts to 4,405.5 mm/year, which occurred in 2017. At the same time, the minimum  
 208 PRCPTOT is found in SM. Iskandar of 1,702.3 mm/year, which occurred in 1997. Furthermore, South  
 209 Kalimantan's annual precipitation variation (PRCPTOT) can be seen in Figure 5(c). The maximum  
 210 PRCPTOT is found in SM. Gusti Syamsir Alam of 4,780 mm/year, which occurred in 1988. Then, for  
 211 the minimum PRCPTOT found in SM. Gusti Syamsir Alam amounted to 1,309 mm/year in 1997.

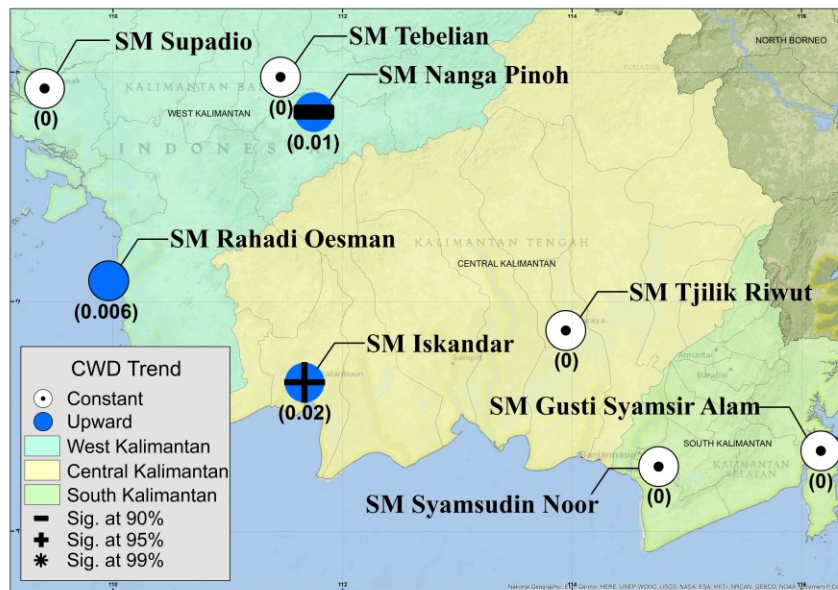


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213 Figure 5. Time series of total precipitation from January 1985 to December 2022 observed at a) West  
 214 Kalimantan, b) Central Kalimantan, c) South Kalimantan



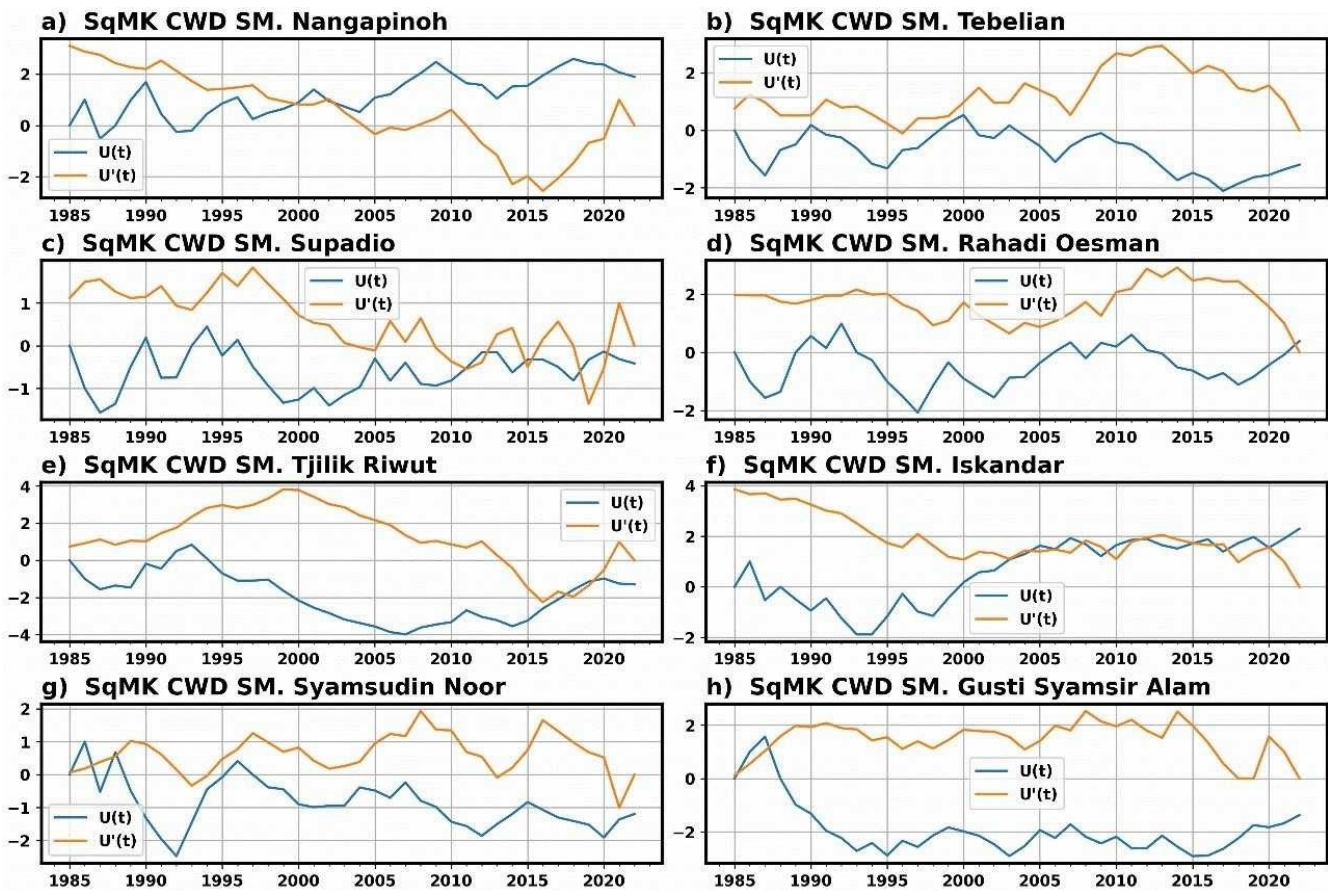
215 The magnitude of PRCPTOT is related to the CWD index. This means that when PRCPTOT is  
216 high, then during one year, the CWD value is also potentially high. Figure 6 shows the trend of CWD  
217 at eight stations. Two stations in West Kalimantan experienced increasing trends, namely SM. Rahadi  
218 Oesman (0.006) and SM. Nangapinoah (0.01). One station was in Central Kalimantan, while the other  
219 five stations experienced no change. The resulting trend of 0 can be seen from Figure 6. A significant  
220 increase occurred in SM. Nangapinoah and SM. Iskandar has significance levels of 90% and 95%,  
221 respectively. Figure 7 shows the changes in the CWD time series in SM. Nangapinoah began in 1999,  
222 while for SM Iskandar, the changes that appeared in the CWD time series started in 2003.



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Figure 6. CWD Trends in West Kalimantan, Central Kalimantan, and South Kalimantan



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Figure 7. Graphical representation of the Mann–Kendall trend test of the CWD time series

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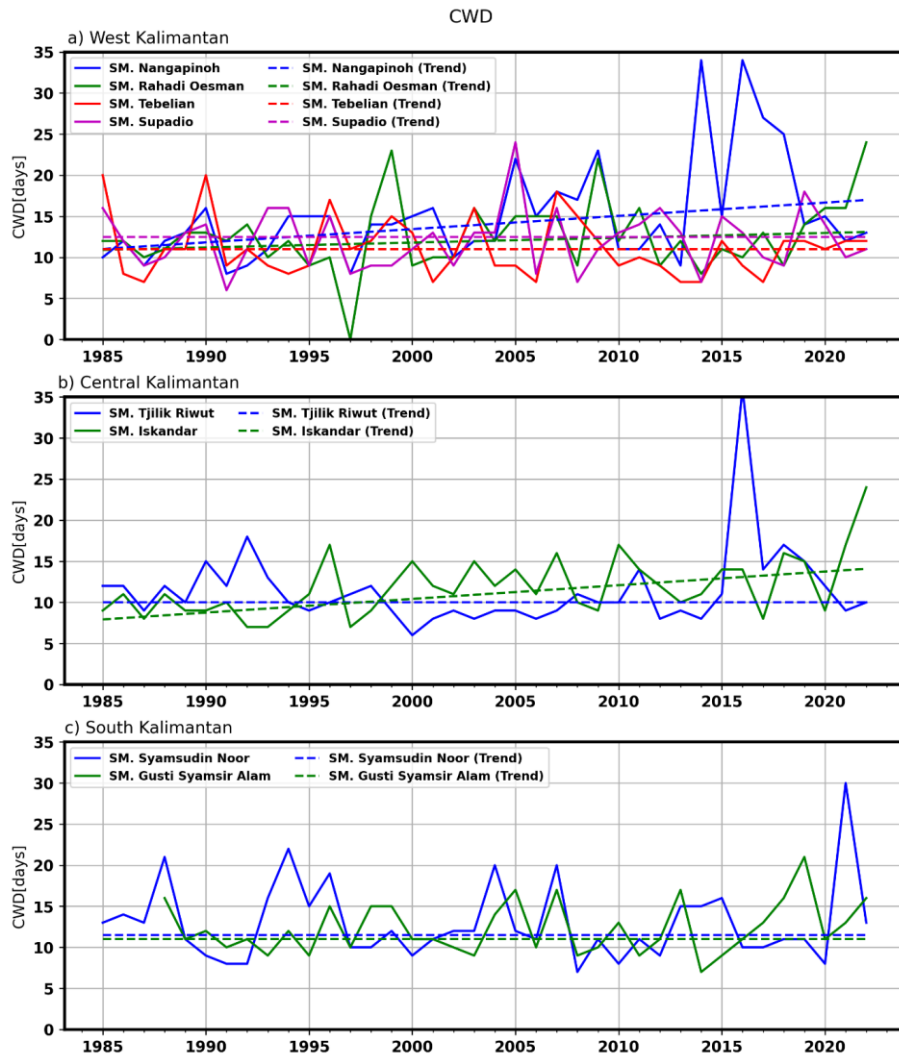
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The variation of CWD over 38 years in West, Central, and South Kalimantan ranges from 6-36 days, as seen in Figure 8. The maximum CWD in West Kalimantan is found in SM. Nangapinoh for 34 days, which occurred in 2014 and 2016 (Figure 8a). Furthermore, in Central Kalimantan, the maximum CWD is found in SM. Tjilik Riwut for 36 days in 2016 (Figure 8b). In South Kalimantan, the maximum CWD value is found in SM. Syamsudin Noor, with the longest rainy day occurred for 30 days in 2021 (Figure 8c). The observations show that the longest rainy season during the 38 years of observation occurred in Central Kalimantan, precisely in the city of Palangkaraya (SM. Iskandar).

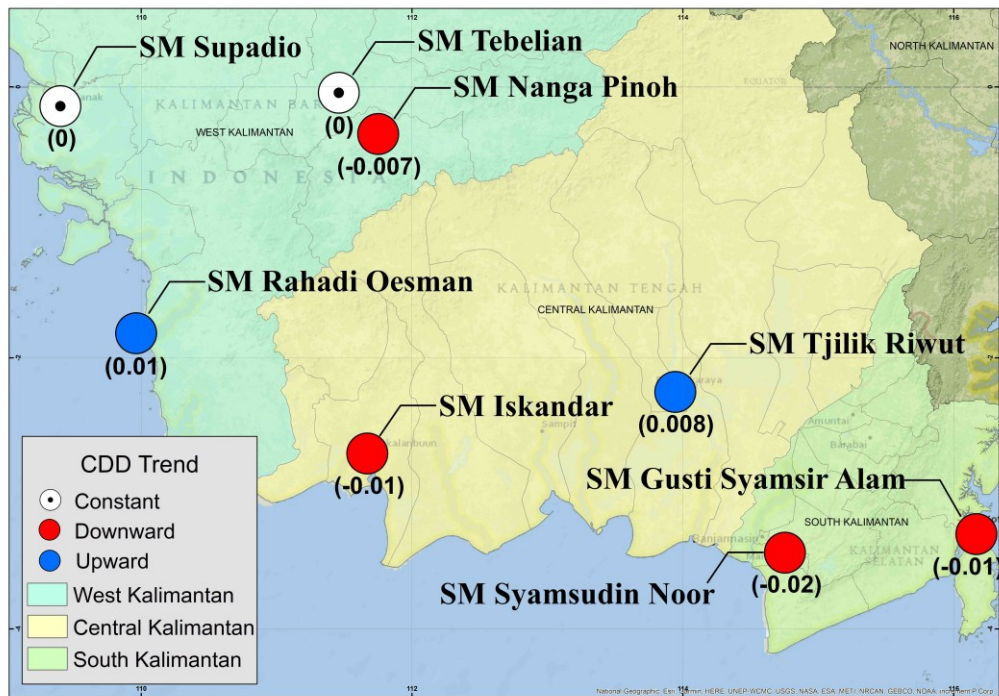


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235 Figure 8. Time series of observed at CWD on a) West Kalimantan, b) Central Kalimantan, c) South  
 236 Kalimantan

237 Figure 9 shows the trend value of CDD at eight observation stations. It can be seen that some  
 238 stations have increased, but some have decreased. The increase in the CDD trend occurred in SM.  
 239 Rahadi Oesman (West Kalimantan) and Tjilik Riwut (Central Kalimantan). Meanwhile, the downward  
 240 trend in CDD occurred in SM. Iskandar (Central Kalimantan), SM. Syamsudin Noor and SM. Gusti  
 241 Syamsir Alam (South Kalimantan). The largest increase in consecutive dry days (CDD) occurred in  
 242 SM. Rahadi Oesman (West Kalimantan) by 0.01 days/year. At the same time, the minimum CDD trend  
 243 is found in SM. Syamsudin Noor (South Kalimantan) by 0.02 days/year.

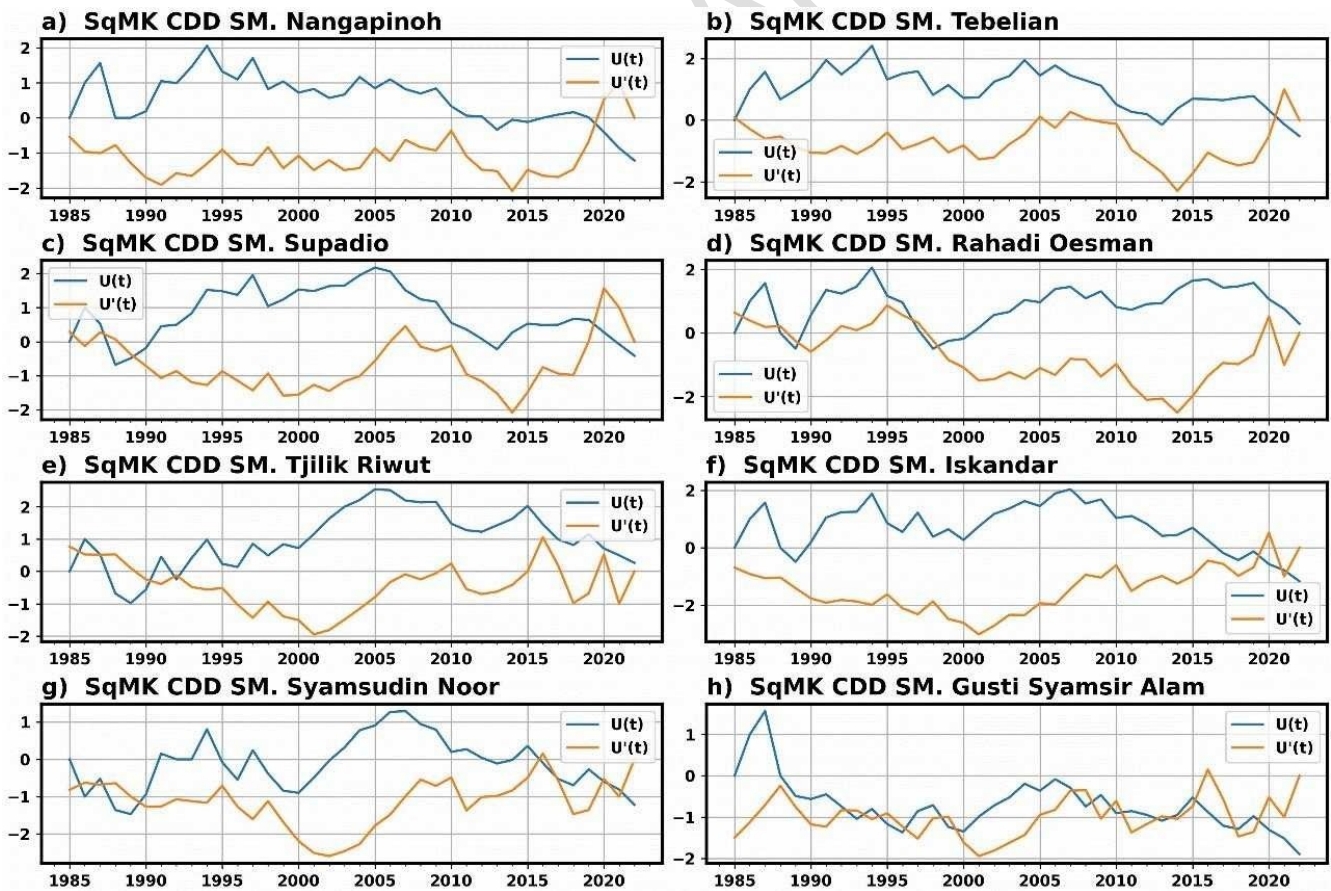




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Figure 9. CDD Trend in West Kalimantan, Central Kalimantan, and South Kalimantan



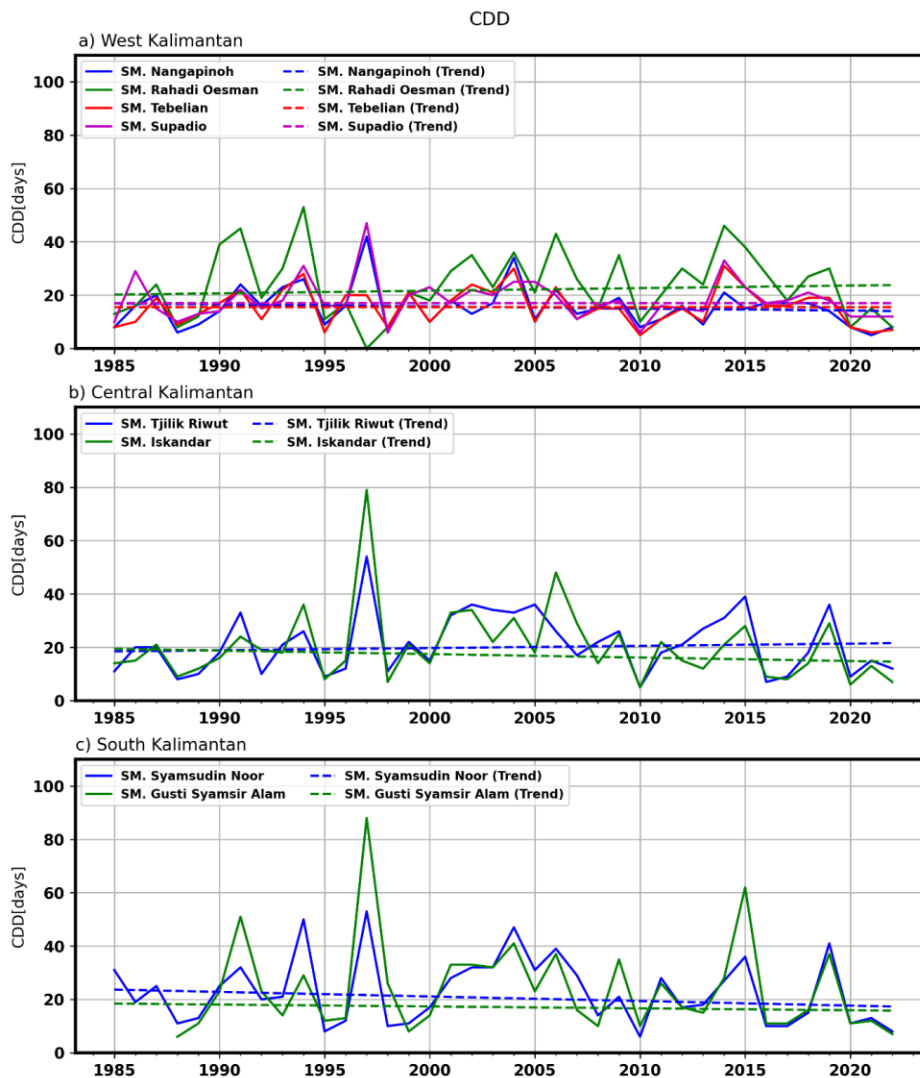
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Figure 10. Graphical representation of the Mann–Kendall trend test of the CDD time series

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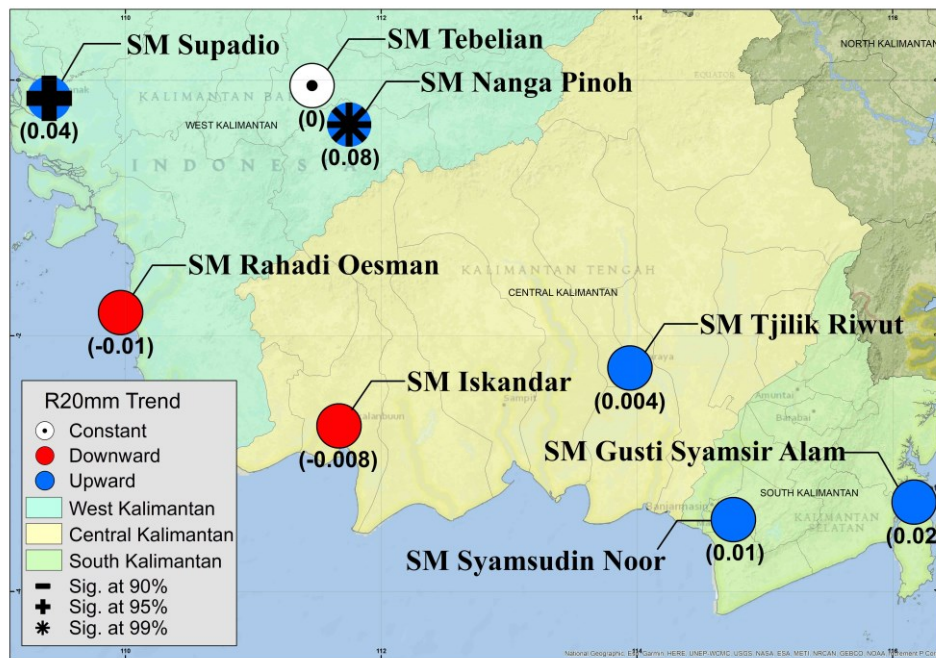
249 The maximum CDD value in West Kalimantan is found in SM. Rahadi Oesman for 53 days,  
250 which occurred in 1994 (Fig. 11a). Furthermore, the maximum CDD value in Central Kalimantan was  
251 found in SM. Iskandar for 79 days SM. Iskandar in 1997 (Fig. 11b). In South Kalimantan, the maximum  
252 CDD value was found in SM. Gusti Syamsir Alam for 88 days, which occurred in 1997 (Fig. 11c). Six  
253 of the eight observation stations showed that the longest number of days without rain occurred in 1997,  
254 where South Kalimantan experienced the longest days without rain compared to West Kalimantan, and  
255 Central Kalimantan.



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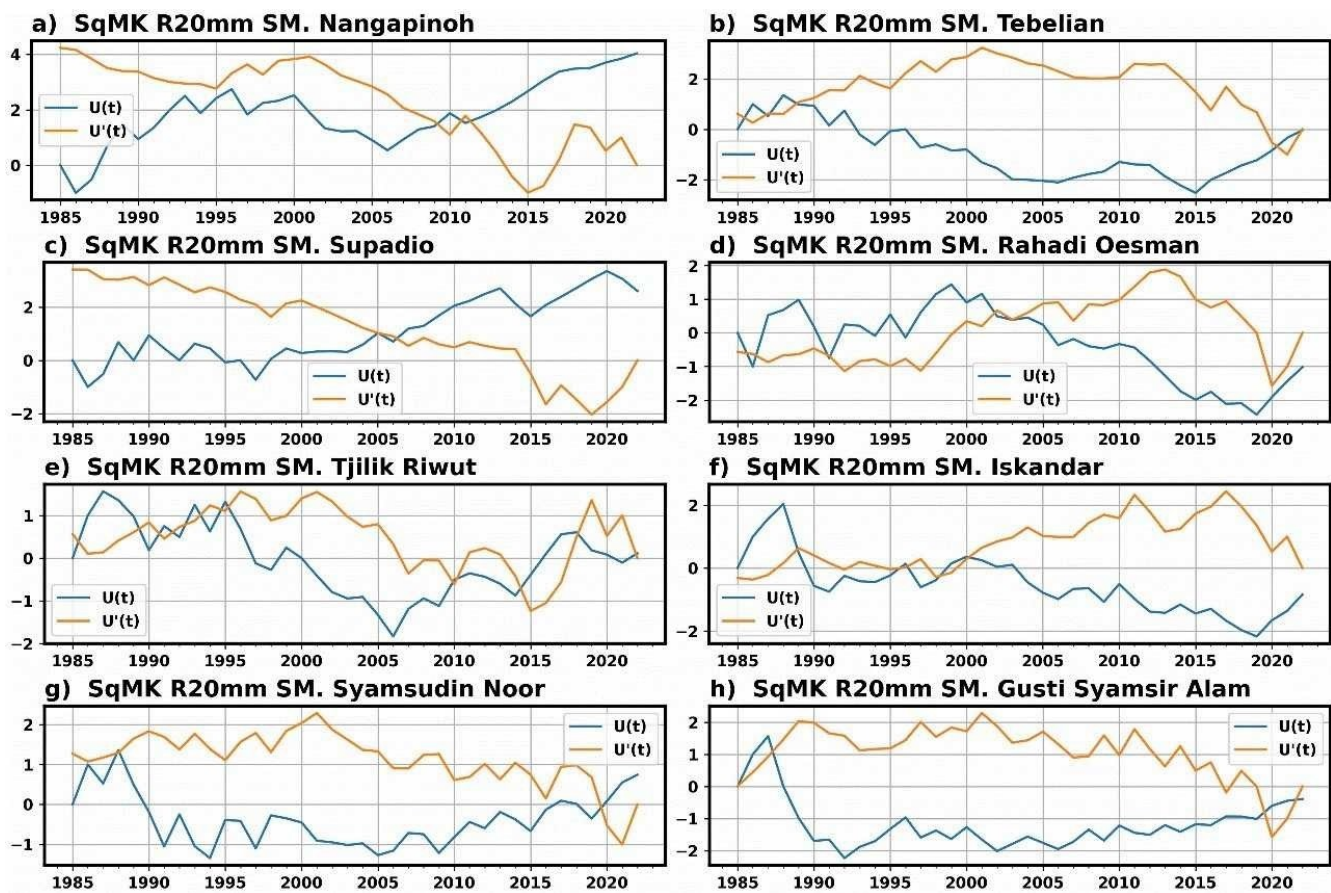
257 Figure 11. Time series of CDD observed at a) West Kalimantan, b) Central Kalimantan, c) South  
258 Kalimantan

259 The R20mm index indicates the number of days with precipitation  $\geq 20$ mm per day. The  
 260 maximum value of this index indicates an area's tendency towards wet conditions. The R20mm trend has  
 261 increased at six stations: SM. Supadio, SM. Nangapinoh, SM. Tjilik Riwut, and SM. Syamsudin Noor,  
 262 and SM. Gusti Syamsir Alam. a significant increase occurred at SM. Supadio and Nangapinoh had  
 263 significance levels of 95% and 99% respectively. Figure 13 shows the changes that appear in the  
 264 R20mm time series at SM. Supadio starting from 2005 while SM. Nangapinoh the changes appear from  
 265 2009. The stations that experienced a decrease in the R20mm trend were SM. Rahadi Oesman and SM.  
 266 Iskandar.



267  
 268 Figure 12. R20mm Trends in West Kalimantan, Central Kalimantan, and South Kalimantan





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Figure 13. Graphical representation of the Mann–Kendall trend test of the R20mm time series

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SM. Nangapinoh (West Kalimantan) obtained the maximum R20mm increasing trend with a

272

significance level of 99%. SM. Rahadi Oesman (West Kalimantan) experienced a decrease in the

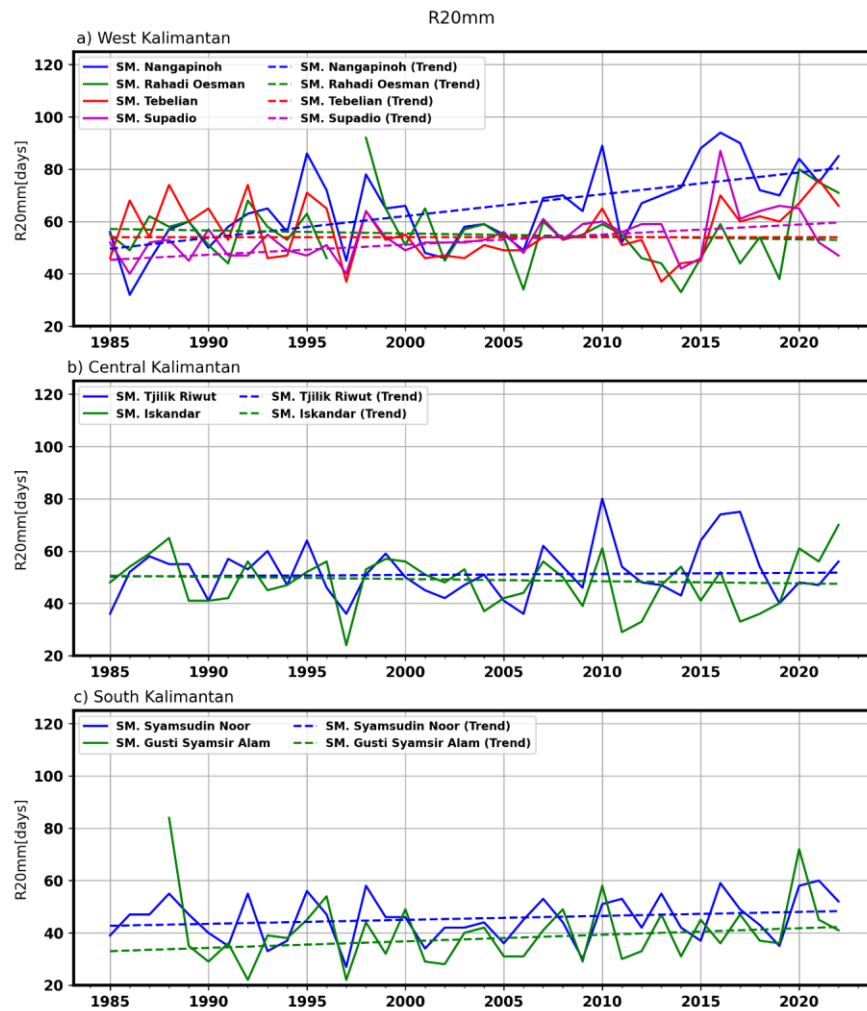
273

number of days with precipitation intensity  $\geq 20\text{mm}$  by 0.01 days per year. Furthermore, the variation

274

of R20mm value in each region can be seen in Figure 14.

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276 Figure 14. Time series of R20mm observed in a) West Kalimantan, b) Central Kalimantan, c) South  
 277 Kalimantan

278 The maximum R20mm value in West Kalimantan occurred in SM. Nangapinoh in 2016, where  
 279 94 days of rain occurred with an intensity  $\geq 20$ mm. Furthermore, in Central Kalimantan, the maximum  
 280 R20mm value occurred in SM. Tjilik Riwut for 80 days in 2010. For South Kalimantan, the maximum  
 281 R20mm occurred in SM. Gusti Syamsir Alam for 84 days in 1988. Of the three regions, the longest  
 282 rainy days with heavy intensity occurred in West Kalimantan, precisely the Malawi district, which  
 283 occurred for 94 days in 2016.

### 284 3.3. Correlation between extreme precipitation index and ENSO phenomenon

285 Correlation analysis was conducted to determine the influence of ENSO events (Niño 3.4 index)

286 on the extreme precipitation index. The results of the correlation analysis of extreme precipitation with  
287 the Niño 3.4 index in the Kalimantan region show positive and negative values. Niño 3.4 correlation  
288 with PRCPTOT, R20mm, and CWD showed negative values. At the same time, the correlation is  
289 positive for CDD and Niño 3.4 index. The largest ENSO influence on PRCPTOT occurs in South  
290 Kalimantan (SM. Syamsudin Noor). This is shown based on the Niño 3.4-PRCPTOT index correlation  
291 value of -0.56 with a significance level of 99%. In addition, there are two stations with significant  
292 correlations of -0.38 (SM. Nangapinoh) and -0.42 (SM. Iskandar) at the 99% significance level in West  
293 Kalimantan and Central Kalimantan. As well as five other stations in West Kalimantan, Central  
294 Kalimantan, and South Kalimantan (SM. Rahadi Oesman, SM. Tebelian, SM. Supadio, SM. Tjilik  
295 Riwut, and SM. Gusti Syamsir Alam) were also significantly correlated at the 95%-98% significance  
296 level. This significant correlation is in line with the results of previous research, which states that  
297 precipitation in Kalimantan has a significant correlation with ENSO (Lestari, dkk., 2018), with the peak  
298 of ENSO events between September-November (Chang, et al., 2003).

299 The CDD index significantly correlated with ENSO events at the 99% significance at all  
300 observation stations (West Kalimantan, Central Kalimantan, and South Kalimantan). For R20mm, six  
301 out of eight stations obtained a significant correlation ranging from 90% to 99% significance level.  
302 ENSO events have a significant correlation with the extreme precipitation index during the SON season  
303 for West, Central, and South Kalimantan. Meanwhile, the CWD index is not significantly correlated  
304 across all observation stations.

### 305 **3.4. Correlation of extreme precipitation index with IOD phenomenon**

306 The correlation analysis of the extreme precipitation index with DMI was conducted to  
307 determine the magnitude of DMI's influence on each extreme precipitation index. In general, the  
308 correlation results show the influence of DMI on extreme precipitation in West Kalimantan, Central  
309 Kalimantan, and South Kalimantan. This is shown based on the correlation value of DMI on the

310 extreme precipitation index. The DMI-CDD correlation is positive (0.03-0.43) at all observation  
311 stations. However, a significant DMI-CDD correlation was only observed at SM. Iskandar (Central  
312 Kalimantan) and SM. Syamsudin Noor (South Kalimantan) had 98% and 99% significance levels,  
313 respectively. Furthermore, for the PRCPTOT index, only SM. Rahadi Oesman (West Kalimantan)  
314 obtained a significant correlation of -0.28 at the 90% level. Furthermore, the R20mm index only has  
315 two stations that obtain a significant correlation, namely SM. Rahadi Oesman (West Kalimantan) and  
316 SM. Iskandar (Central Kalimantan) has a significance level of 95%. Compared to ENSO, IOD has less  
317 influence on the extreme precipitation index, as indicated by the small number of stations that obtained  
318 a significant correlation. Meanwhile, CWD did not obtain a significant correlation with DMI for all  
319 observation stations.

#### 320 **4. Discussion and Conclusion**

321 Analysis of extreme precipitation trends in the West, Central, and South Kalimantan regions  
322 from 1985-2022 shows positive and negative trends. The positive trend of PRCPTOT at eight  
323 observation stations shows that for 38 years, the Kalimantan region has tended to be in wet  
324 conditions. This result has also shown by As-syakur et al., (2013). The largest and most significant trend  
325 increase occurred in SM. Nangapinoh (West Kalimantan). This PRCPTOT trend is correlated to ENSO  
326 by (-0.38) with a significance level of 99%. This aligns with the results of previous research conducted  
327 by Nguyen-Thanh et al. (2023) from 1979 to 2019. The maximum PRCPTOT in West Kalimantan  
328 occurred in 2016 at 5781.7 mm/year. This is because in that year, there was a negative IOD  
329 phenomenon (Fannia et al., 2021). In addition to PRCPTOT, the average total annual precipitation for  
330 West Kalimantan is also greater than that of Central and South Kalimantan. This indicates that West  
331 Kalimantan tends to be wetter than Central and South Kalimantan. Among the three regions, South  
332 Kalimantan is the region with the lowest average total annual precipitation. This result is also shown by  
333 the research of Sukmara et al., (2022) and Supari et al., (2016). Furthermore, in Central Kalimantan,

334 the correlation between PRCPTOT and ENSO is -0.42 (SM. Iskandar) and -0.39 (SM. Iskandar)  
335 significant at the 99% and 98% levels. South Kalimantan, the correlation between PRCPTOT and  
336 ENSO is at SM. Syamsudin Noor and SM. Syamsir Gusti Alam was -0.565 and -0.612, respectively,  
337 with a significance level of 99% and 98%.

338 Compared to PRCPTOT, the increasing trend of CWD is less, ranging from 0.006-0.02 days per  
339 year. This is because ENSO and IOD have a low correlation ( $r \leq 0.24$ ) with CWD in West Kalimantan,  
340 Central Kalimantan, and South Kalimantan. For 38 years, only three stations had an increase in the  
341 trend of CWD, while the rest did not change. Research conducted by Ramadhan et.al (2022) in the East  
342 Kalimantan region also showed an increase in CWD during 2001-2020. An increase in the trend of  
343 CWD occurred at SM. Iskandar (Central Kalimantan), SM. Nangapinoh and Rahadi Oesman (West  
344 Kalimantan). In 2014 and 2016, West Kalimantan experienced the longest rainy days (CWD) of 34  
345 days, especially in SM. Nangapinoh. The maximum CWD in 2016 was caused by the negative IOD  
346 phenomenon during JJA and SO (Fannia et al., 2021). The same results also occurred in Central  
347 Kalimantan. This region had 36 consecutive rainy days in 2016, precisely in SM. Tjilik Riwut.  
348 Meanwhile, in South Kalimantan, the maximum CWD (30 days) in 2021 occurred in SM. Syamsudin  
349 Noor. The event is the impact of the La Niña phenomenon (Novianti et al., 2023).

350 This is not the case with the CWD index. CDD showed increasing and decreasing trends in  
351 West, Central, and South Kalimantan but no significant trends as shown by Supari et al., (2016). SM.  
352 Rahadi Oesman (West Kalimantan) and Tjilik Riwut (Central Kalimantan) obtained increasing trends,  
353 SM. Nangapinoh (West Kalimantan), SM Syamsudin Noor and SM. Gusti Syamsir Alam (South  
354 Kalimantan) had a decreasing trend and SM. Tebelian and Supadio had no change in trend. A  
355 significant correlation between CDD and ENSO occurred at all observation stations (West Kalimantan,  
356 Central Kalimantan, and South Kalimantan) with a significance level of 99%. The results showed that  
357 the maximum CDD in the West Kalimantan region was in SM. Rahadi Oesman for 53 days precisely in



358 1994. This long dry season was caused by the positive IOD phenomenon (Lestari et al., 2018).  
359 Meanwhile, the dry season in Central Kalimantan is longer than in West Kalimantan. In 1997, there  
360 were 79 consecutive days without rain in SM. Iskandar. This event was caused by the El Niño  
361 phenomenon and positive IOD in 1997 (Lestari et al., 2018). The same thing happened in South  
362 Kalimantan for 88 days in 1997, especially in SM. Gusti Syamsir Alam. In line with the research results  
363 by Lestari et al. (2019), the frequency of extreme precipitation strongly correlates with ENSO and IOD  
364 in the dry season. Increasing and decreasing trends in CDD were also experienced in central and east  
365 central Asia from 1981 to 2005 (Rai et al., 2024).

366 Compared to PRCPTOT, CDD, and CWD, the R20mm trends in West, Central, and South  
367 Kalimantan show more varied values. The trends obtained are increasing, decreasing, and constant.  
368 Five of the eight stations tended to be wet during the 38 years of observation. This result has also  
369 shown by Supari et al., (2016). The West Kalimantan region, especially SM. Nangapinoh and Supadio  
370 showed significant increasing trends at 99% and 95%. SM. Rahadi Oesman showed a decreasing trend  
371 and SM. Tebelian had no change. Meanwhile, in the Central Kalimantan region, an increasing trend  
372 was seen in SM. Tjilik Riwut. SM. South Kalimantan that obtained an increase in trend were SM.  
373 Syamsudin Noor and Gusti Syamsir Alam. The maximum R20mm (94 days) occurred in the West  
374 Kalimantan region in SM. Nangapinoh in 2016. This is because 2016, there was a negative IOD  
375 (Fannia et al., 2021). Central Kalimantan region experienced heavy rains for 80 days in 2010,  
376 especially SM. Tjilik Riwut. 2010, it coincided with the La Niña phenomenon La Niña phenomenon  
377 (Lestari et al., 2018). Study Patle & Libang (2014) stated that the increasing trend of precipitation has  
378 an impact on flooding. The increasing trend of R20mm that occurred in West Kalimantan was also  
379 experienced by the central Asian region during 1981-2005 (Rai et al., 2024).

380 Extreme precipitation trends in the West, Central, and South Kalimantan regions have increased  
381 over 38 years. PRCPTOT trends have increased at all observation stations. A significant increase in

382 PRCPTOT trends occurred at SM. Nangapinoh and SM. Supadio precisely in the West Kalimantan  
383 region. Meanwhile, CWD also significantly increased at two stations, namely SM. Nangapinoh (West  
384 Kalimantan) and SM. Iskandar (Central Kalimantan) have significance levels of 90% and 95%. The  
385 number of rainy days with intensity  $\geq 20$ mm significantly increased at SM. Nangapinoh and Supadio  
386 have significance levels of 99% and 95%. In addition, most of the extreme precipitation indices are  
387 significantly correlated with ENSO. The same research results were also shown in the Central Java  
388 region with observation time from 1990-2019 (Firmansyah et al., 2022), North Sumatra in the period  
389 1981-2010 (Irwandi et al., 2018).

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